



University College Dublin
An Coláiste Ollscoile, Baile Átha Cliath

WINTER EXAMINATIONS 2005

SCMXF0028
SCMXP0028
M.Sc. in Meteorology

Dynamical Meteorology
MAPH P310

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Time Allowed: 3 hours

Instructions for Candidates

Full marks will be awarded for complete answers
to **four** questions.

All questions carry equal marks.

Please do not use red pen on the answer books.

Please use separate answer book for each question.

Instructions for Invigilators

Non-programmable calculators may be used during this examination.

Question 1.

a) Let \mathbf{A} be an arbitrary vector whose Cartesian components in a frame of reference rotating with angular velocity $\bar{\Omega}$ are

$$\mathbf{A} = iA_x + jA_y + kA_z \quad (1.1)$$

Write down the relationship between the total derivative of \mathbf{A} in an inertial reference frame ($D_a\mathbf{A}/Dt$) and the corresponding total derivative in the rotating system ($D\mathbf{A}/Dt$).
b) Show that

$$\frac{D_a \bar{i}}{Dt} = \bar{\Omega} \times \bar{i} \quad (1.2)$$

c) Assuming that expressions similar to (1.2) hold for the total derivatives of \mathbf{j} and \mathbf{k} in the inertial frame, what is the resulting relationship between ($D_a\mathbf{A}/Dt$) and ($D\mathbf{A}/Dt$)?

Question 2.

a) Using the Lagrangian control volume method, derive the continuity equation

$$\frac{1}{\rho} \frac{D\rho}{Dt} + \nabla \cdot \bar{U} = 0 \quad (2.1)$$

in (x,y,z,t) coordinates, where ρ is the density and \mathbf{U} is the 3D velocity vector.

b) Again using the Lagrangian control volume method, derive the continuity equation in isobaric coordinates for a hydrostatic atmosphere. Comment on the advantages of this equation.

Question 3.

- a) State the assumptions used in constructing the shallow water model with a rigid horizontal upper lid, bottom topography $h_s(x,y)$ and fluid depth $h(x,y)$ on a β -plane.
b) Starting from the primitive equations of motion on a β -plane and using the above assumptions, show that the governing equations for the model in question are

$$\frac{Du}{Dt} = fv - \frac{\partial \Phi_T}{\partial x} \quad (3.1)$$

$$\frac{Dv}{Dt} = -fu - \frac{\partial \Phi_T}{\partial y} \quad (3.2)$$

$$\frac{Dh}{Dt} = -h \left(\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} \right) \quad (3.3)$$

where $\Phi_T = p_T/\rho$, p_T being the pressure at the upper lid and ρ being the density.

c) Using (3.1), (3.2) and (3.3) derive the potential vorticity equation

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$$\frac{D}{Dt} \left[\frac{\zeta + f}{h} \right] = 0 \quad (3.4)$$

for this model (ζ = relative vorticity).

Question 4.

- a) The vorticity equation for the shallow water model with a rigid lid and a flat bottom on a β -plane can be written

$$\frac{D}{Dt} [\zeta + f] = 0 \quad (4.1)$$

where ζ is the relative vorticity and f is the Coriolis parameter. Show that in terms of the streamfunction ψ eq. (4.1) can be written

$$\frac{\partial}{\partial t} (\nabla^2 \psi) + J(\psi, \nabla^2 \psi) + \beta \frac{\partial \psi}{\partial x} = 0 \quad (4.2)$$

where ∇^2 is the horizontal Laplacian and J is the Jacobian.

- b) Using (4.2) show that the frequency of 2D Rossby waves superimposed on a basic current \bar{u} is

$$\nu = \bar{u}k - \frac{\beta k}{K^2} \quad (4.3)$$

where $K^2 = k^2 + l^2$ and (k, l) are the wavenumbers in (x, y) .

Hint: assume $\psi = \bar{\psi} + \psi'(x, y, t)$ and linearize.

- c) Calculate the group velocity for the waves described by (4.3).

Question 5.

- a) The quasi-geostrophic potential vorticity equation for the free-surface shallow water model on a β -plane ($f = f_0 + \beta y$) is given by

$$\frac{D_g}{Dt} \left[\nabla^2 \psi + f - \frac{\psi}{\lambda^2} \right] = -\frac{f_0}{\Phi_0} \frac{D_g \Phi'_s}{Dt} \quad (5.1)$$

where D_g / Dt is the material derivative following the geostrophic motion, ψ is the geostrophic streamfunction, ∇^2 is the horizontal Laplacian, $\lambda = \sqrt{\Phi_0} / f_0$, Φ_0 is the mean geopotential of the free surface and Φ'_s is the geopotential of the orography (assumed small by comparison with Φ_0).

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Using the linearized perturbation form of (5.1) for a resting basic state with $\Phi'_s = 0$, derive the phase speed for a 1D Rossby wave propagating in the x-direction in this model.

- b) Again using (5.1) but neglecting β , derive the solution for the perturbation streamfunction ψ' in the case of steady motion forced by a mean current \bar{u} blowing over orography of the form

$$\Phi'_s = \text{Re}[\hat{\Phi}_s \exp(ikx)] \quad (5.2)$$

Hint: Since the forcing has spatial dependence of the form $\exp(ikx)$, the solution may also be assumed to have spatial dependence of this form.

Question 6.

- a) Given the following governing equations for the two-layer model of baroclinic instability

$$\left(\frac{\partial}{\partial t} + U_1 \frac{\partial}{\partial x}\right) \frac{\partial^2 \psi_1'}{\partial x^2} + \beta \frac{\partial \psi_1'}{\partial x} = \frac{f_0}{\delta p} \omega_2' \quad (6.1)$$

$$\left(\frac{\partial}{\partial t} + U_3 \frac{\partial}{\partial x}\right) \frac{\partial^2 \psi_3'}{\partial x^2} + \beta \frac{\partial \psi_3'}{\partial x} = -\frac{f_0}{\delta p} \omega_2' \quad (6.2)$$

$$\left(\frac{\partial}{\partial t} + U_m \frac{\partial}{\partial x}\right) (\psi_1' - \psi_3') - U_T \frac{\partial}{\partial x} (\psi_1' + \psi_3') = \frac{\sigma \delta p}{f_0} \omega_2' \quad (6.3)$$

where $U_m = (U_1 + U_3)/2$ and $U_T = (U_1 - U_3)/2$, show that the rate of change of the sum of the kinetic and available potential energies is

$$\frac{d}{dt} (K' + P') = 4\lambda^2 U_T \overline{\psi_T \frac{\partial \psi_m}{\partial x}} \quad (6.4)$$

where $\psi_m = (\psi_1' + \psi_3')/2$, $\psi_T = (\psi_1' - \psi_3')/2$, $\lambda^2 = (f_0)^2 / [\sigma(\delta p)^2]$ and $\overline{(\)}$ denotes an average over the wavelength of the disturbance.

- b) Discuss the energetics of baroclinic waves qualitatively using a box diagram involving \bar{P} (the mean available potential energy), P' and K' .
c) Discuss the physical mechanism of baroclinic instability in terms of the slope of particle trajectories ('the wedge of instability').

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